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Assessment of drainage network analysis methods to rank sediment yield hotspots

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Abstract

This paper aims to test different methods used for assessing the indices of sediment yield to

identify hotspots and rank sediment yield hotspots. This process includes the assessment of

the Entropy Weighting (EW), Fractal Dimension (FD), Slope Length (SL) gradient, and

Sediment Connectivity (SC) methods. The indices at different sub-catchment levels were

applied in the Ilanlu catchment (Iran) and organized based on five different levels of sediment

hazard classes. To assess the performance of sediment hazard mapping, the superimposing

methods were used and assessed by the Erosion Potential Model (EPM). The superimposing

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method showed that 8, 10, 4 and 9 sub-catchments based on the degree of susceptibility obtain the highest results considering the results of FD, EW, SL and SC with an output of EPM model, respectively. The results show that EW and SC can achieve greater performance than FD and SL methods in identifying sediment production hotspots.

Keywords: Cluster analysis; sediment yield; modelling; soil erosion.

1. Introduction

Soil erosion and sediment-related problems are associated with dramatic adverse impacts on the environment and agriculture including economic loses, water quality, reducing the storage capacity of reservoirs and agricultural production (Alexakis *et al.* 2013; Nasre *et al.* 2013; Shit *et al.* 2015). Due to the huge cost of watershed management actions for reducing the consequences of soil erosion and sediment-related problems (Kiani-Harchegani *et al.* 2018; Kiani-Harchegani and Sadeghi, 2020), the areas vulnerable to erosion should be identified to justify and develop sustainable and clean soil and water conservation measures. To this end, many algorithms have been developed to prioritize at the sub-catchment scale based on sediment control measurements to helps watershed managers to manage easily and cost-effectively (Kouli *et al.* 2009; Kalantari *et al.* 2019).

Among catchment geomorphology characteristics, the river network is known as one of the most important impacts on soil erosion, flood, and sediment yield processes. This is because of its implication on water quality, quantity and potential uses at different scales (Al-Dulaimi and Younes, 2017; García-Ruiz *et al.* 2017; Yegemova *et al.* 2018). Therefore, an accurate description of river network geomorphology is crucial for investigating the river flow, sediment transport and flood (Ariza-Villaverde *et al.* 2013; Zhang *et al.* 2015).

The methods of description of the river network were divided into regular and irregular groups by several researchers in the past. In the former group, only the flow pattern

characteristics, such as drainage network length and drainage network gradient, are used to describe the drainage network (Mandelbrot, 1967; Snow, 1989; Prosser *et al.* 2001; Snelder and Biggs, 2002; Piégay *et al.* 2005; Zhang *et al.* 2015). The sediment connectivity (SC) and slope length gradient (SL) are two well-known regular methods employed for the analysis of different soil erosion and sediment dynamics processes. (e.g. Marçal *et al.* 2017; Turnbull *et al.* 2019; Poeppl *et al.* 2019; Najafi *et al.*, 2020).

Llena *et al.* (2019) used the SC method to assess the effects of decadal-scale land use and topographic changes on the transfer of water and sediment through coupling relationships among its components. In another application, Troiani *et al.* (2008) used the SL method as a proxy of stream power to delineate the capacity of the drainage network to transport sediment and erode its bed.

Although the regular methods are widely used, results of studies by Hassan and Kurths (2002), Zhang *et al.* (2015), and Malekinezhad *et al.* (2017) suggest that these regular methods may not fully represent the current drainage networks irregularities, emphasizing the need for alternative analyses. As a result, there is a growing use of methods, which are relatively newer than the SC and SL, for analysis of drainage networks. These methods have been developed based on irregular properties of objects (i.e. drainage network) and provide more fitting descriptions of these characteristics than regular methods (Mandelbrot, 1967; Snow, 1989; Zhang *et al.* 2015; Sepehri *et al.* 2018). The fractal dimension (FD) is also one of the most popular methods in this category. Recent applications of FD are not only focused on geomorphological studies (Nur *et al.* 2016; Hu *et al.* 2020) but also on other topics such as astronomy (De Cola and Lam 2002; Kusák, 2014), geology (Zhang *et al.* 2020; Dou *et al.* 2020), meteorology (Zhou *et al.* 2020; Yao *et al.* 2020) and hydrology (Zhang *et al.* 2015, Al-Wagdany *et al.* 2020).

Among others, entropy weighting (EW) is another method utilized to assess the drainage network irregularity, precipitation, water quality, soil moisture, groundwater networks, and sediment studies (Su and You, 2014; Stosic *et al.* 2017; Keum *et al.* 2017; Pournader *et al.* 2018; Arabameri *et al.* 2019). Depending on which property of the drainage network is considered, the EW can be categorized as irregular or regular method. For example, Fiorentino *et al.* 1993 employed a regular method "entropy of the Horton-Strahler ordering" for drainage network analysis. The EW method will be classified as an irregular method if it is used to assess drainage network irregularities and relationship with other aspects of studies on natural hazards, such as soil erosion, sediment yields or flood.

Based on best our knowledge, on one hand, there is scarce information about applications of EW for assessing the irregularity property of drainage network and its relationship with sediment yield, flood hazard and other natural hazard studies. On the other hand, it is very limited the comparison studies between irregularity and regularity properties of drainage network related to hydrology studies in developing countries with catchments that registered high sedimentation problems. Therefore, to fill this gap, this research aims to compare regular and irregular methods to prioritize sub-catchments by sediment yield. This process includes the assessment of the Entropy Weighting (EW), Fractal Dimension (FD), Slope Length (SL) gradient, and Sediment Connectivity (SC) methods. These indices at different sub-catchment levels were first estimated and then organized based on five different levels of sediment hazard classes using hierarchical clustering. To assess the performance of sediment hazard mapping, the superimposing methods were used and the modelling results were assessed by EPM (i.e. Erosion Potential Model) method. To achieve this goal, we selected the Ilanlu catchment in the Northwest Hamadan, Iran.

2. Materials and Methods

2.1. Study area

The study site is located in a catchment with an area of 17.89 km² in the northwest of Hamadan Province, Iran. According to Asad-Abad Weather Station's data (31°24'45" to 31°27' 29" north; 41°55'20" to 41°57'34" east) collected between 1997 and 2015, near the Ilanlu catchment, the average temperature of the study site is 10.8 (-15 to + 34). February and August are the coldest and hottest months, respectively. The annual average precipitation is 443 mm. According to the ombrothermic curve, the driest months of the year are May to September, which corresponds to a semi-arid region. The predominant drainage network pattern is dendritic at some parts of the catchment, which becomes parallel due to the hillslope steepness, which is characterized by the presence of lime covered by a clay layer (Fig. 1). In the upstream catchment sections, a small portion of precipitation reaches lower layers. This amount of water flows from springs and the remaining portion of the rainfall runs off over the catchment where the permeable lithic-rich facies intersect non-permeable sediments (Ildoromi et al. 2019).

Fig. 1.

2.2. Methodology

This section briefly explains the drainage network assessment methods under investigation. These methods reveal irregularities of drainage networks and show their relationships with sediment productivity/transport and possible yield.

2.2.1. Regularity Methods

- Stream Length gradient (SL)

This method depends strongly on the flow power (Bizzi and Lerner, 2013; Troiani *et al.* 2017; Moussi *et al.* 2018). In a specific reach of a drainage network, the stream power is a measure of hydraulic driving forces associated with sediment detachment and transport processes (Burbank and Anderson, 2001; Troiani and Della Seta, 2008). Troiani *et al.* (2017) and Yang *et al.* (2019) used the SL method to define a measure of river flow power in a unit of length. They also used this method to define a portion of the river flow energy at a certain reach of the river and it relates to the capability of a stream to transport sediments and to erode its bed (Table 1).

- Sediment Connectivity (SC)

The relationship between flow and sediment, as two major attributes of sediment distribution throughout the catchment, reveals the linkage between the sediment site and upstream discharge into the downstream (Sandercock and Hooke, 2011; Sougnez *et al.* 2011; López-Vicente *et al.* 2013). This is a very popular index among researchers due to simplicity, minimum information requirement, and complementation with field observations. According to Estrany et al. (2019), López-Vicente *et al.* (2015), Sougnez *et al.* (2011), Zhang *et al.* (2017) and Mahoney *et al.* (2018), the high SC value indicates the predominance of erosion over sedimentation and that the desired point typically has a high rate of sediment transportation of sediment continuity (Table 1).

2.2.2. Irregularity Methods

-Fractal Dimension (FD)

The fractal uses to describe irregular natural objects. Contrary to the Euclidean method used merely to describe regularly objects (e.g. 0 for point objects and 1, 2 and 3 values for lines, Plans and volume objects, respectively) (Chen *et al.* 2019), the fractal theory determines the dimension of irregularity or fractal objects using the self-similar properties, which can be categorized as 1< d < 2 and 2 < d < 3 for two-dimensional and three-dimensional spaces (Table 1). These methods eliminate the barriers of regular ones in one-, two- and three-dimensional spaces. Therefore, the fractal offers a more comprehensive account of the real properties and their irregularity (Zhang *et al.* 2015; Malekinezhad *et al.* 2017). Although self-similar structure should not necessarily be fractal, fractals typically involve self-similarity (Zhang *et al.* 2015; Malekinezhad *et al.* 2017). In the field of hydrology, the FD is used to present the drainage network irregularity and its relationship with soil erosion, flood and sediment yield (Ariza-Villaverde *et al.* 2013).

- Entropy Weighting (EW)

EW measures the properties of such structures as disorder, instability, and uncertainty. First, Stephan Boltzmann (1877) proposed this method and later it was quantitatively described by Shannon (1948). This study used the EW to assess the change in the self-similar properties of the objects (Jaafari *et al.* 2014). The FD uses the regression relationship for such assessments (Table 1).

2.2.3 Workflow

Fig. 2 presents a workflow used to determine and develop various catchment layers to assess the efficiency of different sediment distribution methods across the catchment. To this end, the catchment under investigation was assessed using ArcHydro (as Add-in in ArcGIS 10.7 software) and the automatic catchment delineation tools, which divided into 28 subcatchments the watershed using the topographic map (1:50000 and 20 m of contour interval). Further, the drainage networks of the sub-catchments were regarded as input data for FD and EW to assess irregularity of them and to establish their linkage to the degree of sediment transportation. The SL and SC methods need more than drainage data, including macroroughness and slope layer data obtained from applying DEM to the catchment (Fig. 2). After collecting the required data and measuring the susceptibility of sub-catchments to sediment yield using the aforementioned methods and EPM method, the hierarchical cluster analysis (HCA) was used to classify the total area into five classes, namely very high, high, moderate, low, and very low sediment yield susceptibility.

Table 1.

Fig. 2

2.2.4. The algorithm of clustering to apply SL and SC methods

The FD and EW methods produce only one value of the sediment yield index for each sub-catchment; whereas, the SL and SC methods for every sub-catchment produce a series values of SL and SC indices across river network. In this regard, for assigning a single value of SL and SC indices to every sub-catchments. It is necessary to transform the series values to raster map based on an interpolation method. In this study, the Inverse Distance Weighting (IDW) Interpolation was used to produce the raster map of SL/SC index. Then, the blow algorithm was applied to assign a single value related to SL/SC index for each sub-catchments:

- **Step 1:** Reclassify each sub-catchment into five different classes based on the values of SL or SC indices.
- Step 2: Assigning specific values (w_{ij}) to each classes using AHP (Analytic hierarchy process) (Table 2) and, then, for H_i to each sub-catchment using Eq. 1.

$$H_{i=1:28} = \sum_{j=1}^{5} w_{ij} \cdot x_{ij} \tag{1}$$

Where X_{ij} stands for the percentage of areas under different classes of SL or SC (Stage 1), i represents the number of sub-catchments, and j denotes the degree of SL and SC values (e.g. j=1 is for very low SL and SC and j=5 is for very high SL and SC).

• **Step 3:** Clustering all sub-catchments into five clusters considering the sediment yield rate using the Hierarchical Cluster Analysis (HCA).

Table 2.

2.3. Hierarchical Cluster Analysis (HCA)

After calculating the sediment yield rate for every sub-catchment, the HCA was utilized to identify the sub-catchments that are similar in the sediment yield rate. This technique involves three steps (Ouarda *et al.* 2008).

- 1. Determining the similarity among sub-catchments using a paired-comparison. At this stage, the degree of proximity and similarity is determined using the distance criterion or the difference between the values of the erodibility index among sub-catchments.
- 2. Classification of sub-catchments in a binary hierarchical tree using a linkage function. Here, the linkage function, the paired-sub-catchments with close erodibility scores are placed at the same cluster (step 1). In the next step, all clusters are grouped in a larger cluster. This process continues to form a hierarchical tree.
- 3. Identification of clusters: during this last step, grouping clusters in a hierarchical tree and cutting the hierarchical tree at an arbitrary point are applied for distinguishing the clusters.

To apply the HCA, two methods, namely Minkowski distance (Equation 2) and Ward's linkage (Equations 3 and 4) (Lin and Chen, 2006; Ouarda *et al.* 2008), as distance and linkage

functions were applied. The Minkowski distance is a metric in a normed vector space, used to measure the distance between two variables. The Ward's method is a sum of Euclidean distances between cluster means for all inter-cluster variables, in that each variable is made incremental. Therefore, an increase in the total within-cluster sum of squares (WSS) leads to the combination of two clusters of p and q (Lin and Chen, 2006; Ouarda *et al.* 2008). The WSS is the total distance between inter-cluster objects and the cluster's centre of gravity (i.e., the mean vector):

$$d(a,b) = \left(\sum_{i=1}^{n} |ai - bi|^{\lambda}\right)^{1/\lambda} \tag{2}$$

where a_i and b_i are the i^{th} arguments of the sets (sediment production rate of the subcatchments); a, b and λ are a parameter such that λ (- ∞ , ∞), which is assumed to be 3

$$WSS_p = \sum_{i=1}^{n_p} d^2(X_{pi}, X_p^-)$$
 (3)

where n_p and x_p stand, respectively, for the size and the centroid of Cluster p. Following this equation, it can be obtained the distance between Clusters p and q:

$$d(p,q) = WSS_{(p+q)} - (WSS_p + WSS_q) = \frac{n_p n_q d^2(X_p, X_q)}{n_p + n_q}$$
(4)

2.4. Evaluation of the performance

The observed sediment data is the necessary condition for using physical-based distributed models to simulate and validate the sediment yield of a basin (Aga *et al.* 2020). One of the most important challenges, especially in developing countries, is related to data-scarce or unavailability of measured sediment data. Therefore, in recent years various empirical models have been extended to predict sediment yield either in the form of suspended or bed load. The acceptable performance of these models has a high dependence on how the required data are measured (Ildoromi *et al.* 2019).

In this study, among the empirical methods, the EPM was used to evaluate the consistency of classified outputs of FD, EW, SC and SL methods. EPM is a well-known empirical method which has been successfully used in many regions, especially in developing countries which data of sediments are scarce (Noori *et al.* 2016; Najafi *et al.* 2020). Therefore, in this study, to estimate the indicators of this method, we used data provided by Department of Natural Resources of Hamadan Province. These data were collected by experts with regular field measurements and surveys. The majority of this data collection was controlled under field conditions, which would reduce the uncertainty of the accuracies of th eoutputs.

The EPM model for calculating erosion (Equations 5 and 6) and sediment yield (Equations 7 and 8) need to six factors i.e. surface geology, soils, topographic features, climatic factors (including mean annual rainfall and mean annual temperature), and land use (Noori *et al.* 2016; Aga *et al.* 2020):

$$Z = Y.X_a(\Theta + \sqrt{I}) \tag{5}$$

$$Wsp = TH\pi Z^{1.5} \tag{6}$$

$$G_{sp} = Wsp. R_u \tag{7}$$

$$Ru = \frac{4(P \times D)^{0.5}}{L + 10} \tag{8}$$

where Z is the erosion coefficient, Y is the coefficient of the rock and soil resistance, Xa is the land use coefficient, Q is coefficient of the observed erosion process, I in per cent is average slope gradient of the sub-catchment, Wsp is the average annual erosion (m³ km² y¹), Ru is sediment yield coefficient, T is the temperature coefficient (°C), H is the mean annual precipitation (mm), P is the perimeter of the sub-catchment (km), L is the length of sub-catchment (km) and D is the difference between mean altitude of the sub-catchment and the altitude of the sub-catchment outlet (km) (Efthimiou and Lykoudi, 2016; Noori $et\ al.\ 2016$). In this study, the value of EPM factors was prepared from General Department of Natural Resources of Hamadan Province (Table 3).

To determine the accuracy of FD, EN, SC and SL methods, the result of EPM model (Fig 2) was superimposed on the output of mentioned methods and then based on the number of common sub-catchments, the percentage of accuracy was accessed.

Table 3

3. Results and Discussion

3.1. Analysis of FD and EW

The FD of each sub-catchment was calculated based on the changes in the counting of boxes (Fig. 3). It is worthy to highlight that box-counting is a method of gathering data for analyzing complex patterns by breaking a dataset, object, image, etc. into smaller and smaller pieces. The results showed that sub-catchment 8 obtained the highest FD (1.14). According to Fig. 1, its drainage network also reached the highest number of branches and irregularities. In contrast, the drainage network of sub-catchment 4, with an FD of 0.89, had the least number of branches and irregularities.

Table 4 shows that the correlation coefficient (R) between FD of a drainage network and box-counting estimate at all sub-catchments is higher than 61%, suggesting that the accuracy of the FD at the river network. These findings are consistent with those of Zhang *et al.* (2015) and Malekinezhad *et al.* (2017), who also used the box-counting method. After obtaining the FD at each sub-catchment, the total area was divided into five classes based on HCA method: very high (1.108-1.136), high (1.088-1.103), moderate (1.037-1.08), low (1.002-1.006), and very low (0.8875-0.9287) (Figs 4a and 5a).

Sub-catchments 2 and 11 show, respectively, the highest (0.06) and lowest (0.008) EW values. In other words, they are the sub-catchments with the highest drainage network irregularities (or higher rate of sediment yield) and lowest drainage network irregularities (or the lowest sediment yield) (Fig. 6). Based on HCA method, this section divided the total area into five classes: very high (0.056-0.065), high (0.039-0.049), moderate (0.026-0.034), low (0.02-0.02), and very low (0.008-0.016) (Figs 4b and 5b).

Table 4

Fig. 3

Fig. 4

Fig. 5

Fig. 6

3.2. Analysis of SL and SC methods

After calculating the SL of river network at all sub-catchments and generalizing it to all sub-catchment surfaces with weighting (IDW)_method, the standard deviation of SL values at those sub-catchments were compared among them. According to these findings, the highest SD (3661.635) was observed in the sub-catchment 4. Also, Sub-catchments 2 and 3 obtained the lowest and highest SL (0 versus 35116.25) among all sub-catchments, respectively (Fig. 7).

The SL internal interval variations of Sub-catchment 4 were used to prepare the classification table. The minimum SL (0) at Sub-catchment 2 and the maximum SL at Sub-catchment 3 (35116.26) were used as boundary conditions in the SL classification table at all sub-catchments. After classifying the sub-catchments into five classes based on SL values, namely very low (0 - 3653.4), low (3653.4 - 6813.11), moderate (6813.11 - 10071.55), high (10071.55 - 14514.89), and very high (14514.89 - 35116.257), a value ranging from 0.04 to 0.52 was assigned to each class at every sub-catchment using AHP method (Table 2). Finally, these ranges were divided into five groups, namely very low (3.75 - 4.40), low (4.53 - 4.9), moderate (5.03 - 5.22), high (5.72 - 6.59), and very high (7.77 - 7.77) sediment yield rate (Figs 5c and 6). Table A1 presents the SL-based clustering algorithm details.

Similar to sub-catchments prioritization with the SC method, the internal values of the classification table were extracted from Sub-catchments 7; also, the highest and lowest values of this table were, respectively, extracted from Sub-catchments 27 and 28 (Fig. 7). After the sub-catchments were divided into five classes, namely very low (-5.23 to -1.84), low (-1.84 to -1.1), moderate (-1.1 to -0.35), high (-0.35 to 0.46), and very high (0.46 to 2.31), based on the SC values, which were assigned to a specific number using AHP method. These numbers were clustered into very low (11.92 - 21.24), low (23.12 - 28.46), moderate (30.23 - 31.87), high (33.69 - 36.94), and very high (40.01 - 40.97) sediment yield susceptibility (Figs 4d and 5d). Table B1 shows the specifications of the SC clustering algorithm. However, it is important to remark that we recognize that at the hillslope scale due to the connectivity processes and different land management or environmental conditions (soils, parent material, vegetation, etc.), these results can vary (Bracken *et al.* 2015; Rodrigo-Comino *et al.* 2018; Cossart *et al.* 2017). It depends on the scale used.

Fig. 7.

3.3. Comparison and Assessment of Accuracy of Irregular and Regular Methods

To compare and assess the accuracy of each sediment yield hotspot and set priority for layers prepared by the above methods, this study used the EPM model.

To evaluate the results of the total sediment (Gs) obtained from the EPM model (Table 3), the volume of the check dams' sediments (Vs) in the sub-catchments (Fig 1), which were collected from the database of the General Department of Natural Resources of Hamadan Province, was used. There is always a difference between observed and estimated sediment data (i.e. Vs and Ss) in nature that can due to the deposition of sediments along the transfer way or their trapping behind vegetation or holes in the catchment (Imeson and Prinsen, 2004; Najafi *et al.* 2020). Therefore, in order to determine the best-performed Ss and Ss curve, different regression fitting procedures were applied. The results of a simple linear regression study showed high correlations were established between them as below equation (9). In this regard, Different statistical criteria, viz., coefficient of determination (Ss), a significant level of the p-value (Ss), root mean square of error (Ss) and coefficient of efficiency (Ss) were then respectively applied to the evaluating the fitness (Sss) and coefficient of efficiency (Ss) in the value of Ss, Ss, Ss, Sss and Sss for equation 9 obtained respectively 0.79, 0.02, 0.72 and 0.007. Therefore, the efficiency of the EPM model in the Ilanlu catchment has been confirmed and can be used to evaluate EW, Sss, and Ssss cmethods.

$$VS = 0.227 \ Gs-223$$
 (9)

In the following, the values of EPM model for each sub-basin such as used methods were classified into five sediment production susceptibility i.e. very high, high, moderate, low and very low based on HCA method (Fig 8). Then, the output of used methods was superimposed to the output of the EPM model. In the FD method, out of 8 sub-catchments are common with the output of the EPM model. In the EW method, 10 sub-catchments are common (Table 5).

Fig. 8

Table 5

These results show the good accuracy of the FD and EW methods in sediment yield priority. A major reason for the superiority of the EW method over the FD method could be the assessment changes in the presence of drainage networks in different box sizes. The regression method in the FD is used to assess this dynamic. Since these changes have a loglog graph, a slight change in linear regression slope causes a large uncertainty in the final sediment yield mapping. For the assessment of this behavior in the EW method, the sum operator is used (Table 1). Therefore, this method is associated with lower uncertainty. The results from the SL method showed that only sub-catchments 4, 5 in classes of very high and high and sub-catchments 10 and 26 are common with EPM model (Table 5). Our results indicated that the sub-catchments 4, 5 and 10 and 26 have the maximum and minimum value of DH/DL or slope degree, respectively. Although this method can demonstrate both stream power and sediment transportation, the inclination is the core factor in this regard (Table 1) as other authors also demonstrated in the mountainous catchment (Brandolini et al. 2016; Nadal-Romero et al. 2014). The SL is typically used in perturbation zones of active geodynamical settings. Therefore, in low perturbation rate areas, the river network has a low SL; whereas, the SL is higher at a river network located in areas with moderate and high subduction rates (El Hamdouni et al. 2008; Troiani et al. 2014; Gaidzik et al. 2017). In SC index, Subcatchment 5 is common with EPM model. This finding shows that this method achieved a 4% accuracy.

In the SC index, there are 9 common sub-catchments with EPM model (Table 5). There are two reasons which could cause this method to have high accuracy. The first reason is the opposite of other methods, is related to considering various parameters such as slope,

roughness coefficient, upslope area and flow length. The second and impotence reason is related to considering the linkage between upstream and downstream areas, which is called the connectivity in the hydrology studies, has not been considered in other methods. Heckmann *et al.* (2015) in their study recommended that in studies of flood, erosion, sediment yield and other hydrology studies if you want to have high accuracy, it is necessary to consider the connectivity indices. Recently, Rodrigo-Comino et al., (2020) also demonstrated that this index could be even applied to smaller scales if *in situ* measurements are taken.

3.4. Challenges and other considerations

Overuse of natural resources can accelerate soil loss and leads to increased sediment loads in streams which is reflected in problems within and beyond individual catchments (Li et al., 2020a, 2020b). Recognition of hotspot areas susceptible to sediment loss is, therefore, an essential element to design effective and comprehensive strategies related to land management at the catchment scale, considering sediment yield and transport (Kiani-Harchegani and Sadeghi, 2020; Najafi et al. 2020). Thus, identifying areas of sedimentary hotspots in management concepts is especially important in developing countries where high erosion and sediment delivery rates cause severe problems but there is a lack of information (Heckmann et al. 2018; Najafi et al. 2018). Therefore, EW and FD as irregular methods and SL and SC as regular ones were evaluated as concepts in assessing the hydrological and erosional responses of watershed components having different natural and management characteristics. The results showed that they can be used as an effective screening tool for hydrological, erosional and sediment management. Such studies would meet the practical need for sediment management at different temporal and spatial scales. We consider that a greater understanding of the above methods and the interaction between soil erosion and

sediment transport processes, as well as their spatiotemporal responses, would decrease uncertainties in interpreting sediment transport and sediment yield within catchments (Najafi *et al.* 2020).

4. Conclusion

This study used four common methods, namely the EW and FD as irregular methods and SL and SC as regular methods, for the prioritization of the sediment production hotspots in the study site. The soil erosion of each sub-catchment was measured using the mentioned methods after dividing the case of study into 28 sub-catchments. The outputs were superimposed to the output results pf EPM model. The EW, followed by the SC, FD, and SL was more accurate in categorizing the sub-catchments from viewpoint of susceptibility to sediment transportation and possible yield, respectively. These results support the superimposing method. Although the results of the current study revealed the superiority of some methods over the others under investigation, each method has its pros and cons. Therefore, it is suggested to improve the performance of these methods or use a combination of them to help developing countries to foresee future and potential hazard risks. On the other hand, these methods are possible to be applied in other drainage networks which have a direct relationship with soil erosion and sediment production. This gives another relevant point to this research.

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Table 1. Summary and examples of the different regular and irregular equations used in Ilanlu catchment

Descriptor	Formula	Reference
Fractal Dimension	$d = \frac{\log(N)}{\log(r_h)}$	Martinez-Mena et al. (1999); Xu et al. (2014); Hui et al.
	$\log(r_b)$	(2017); Sampaio et al. (2019)
Entropy Weighting	$p_{i} = -k \sum_{j=1}^{n} f_{ij} \ln f_{ij}$ $f_{ij} = \frac{r_{ij}}{\sum_{j=1}^{n} r_{ij}}, k = 1/\ln n$ $r_{ij} = \frac{x_{ij} - \min\{x_{ij}\}_{i}}{\max\{x_{ij}\}_{i} - \min\{x_{ij}\}_{i}}$ $\lambda_{i} = \frac{1 - p_{i}}{m - \sum_{i=1}^{m} p_{i}}$	Devkota et al. (2013); Pournader et al. (2018); Chen et al. (2019); Arabameri et al. (2019)
Sediment Connectivity	$SC = log_{10} \left(\frac{D_{up}}{D_{dn}} \right)$ $D_{up} = \overline{WS} \sqrt{A}$ $D_{dn} = \sum \frac{d_i}{\overline{W_i S_i}}$ $\overline{W} = 1 - \frac{(altitude\ mean\ - altitude)}{altitude\ range}$	Sandercock and Hooke (2011); Sougnez <i>et al.</i> (2011); López-Vicente <i>et al.</i> (2013); Sherriff <i>et al.</i> (2019); Kalantari <i>et al.</i> (2019)
Stream Length	$SL = \left(\frac{\Delta H}{\Delta L}\right) \times L$	Hamza et al. (2019); Troiani et al. (2018); Moussi et al.
gradient	$\Delta L = \left(\frac{\Delta L}{\Delta L}\right) \wedge L$	(2018)

d: fractal dimension; N: number of measurement scales; \mathbf{r}_b : small segments of box-counting method (pixel size) \mathbf{r} : length of the phenomenon; \mathbf{r}_{ij} : normalization index; \mathbf{X}_{ij} : number of networks in which the drainage network is available; \mathbf{P}_i : entropy value; λ_i : weight of the ith index; \mathbf{SL} : Stream length gradient; $\Delta \mathbf{H}$: maximum and minimum of elevation of a given stream reach; $\Delta \mathbf{L}$: total length of the segment; \mathbf{L} : distance from the source of the river to a minimum elevation of the segment; \mathbf{SC} : Sediment Connectivity; \mathbf{W} : average weighting factor of the upslope contributing area; \mathbf{A} : (dimensionless and equal to macro roughness); \mathbf{S} : average slope gradient of the upslope contributing area (\mathbf{m}^2); \mathbf{Si} : the slope gradient of the \mathbf{i}^{th} cell (\mathbf{m} \mathbf{m}^{-1}).

Table 2. Assigned weight for the classes of the SL and SC methods

			Primary Score			Weighting	
	Very High	High	Moderate	Low	Very Low	Weighting	
Very High	1	4	5	7	8	0.52	
High	•••	1	4	5	6	0.25	
Moderate			1	4	5	0.13	
Low				1	2	0.05	
Very Low				•••	1	0.04	

Table 3: The values of EPM model parameters and volume of sediment (m³) in chekdams

No. Sub-catchment		Н			P		Wsp	Gsp	A	Gs	VS
	(%)	(m)	L (KM)	D (km)	(km)	Ru	$(m^3km^{-2}y^{-1})$	$(m^3km^{-2}y^{-1})$	(km^{-2})	(m^3y^{-1})	(m^3y^{-1})
1	8.4	2013	1.55	8.46	4.07	2.03	669.36	1359.09	0.45	612.86	168.43
2	9.3	2026	2.29	9.36	6.34	2.51	684.42	1715.71	1.34	2291.78	665.41
3	8.1	2001	1.83	7.33	5.58	2.16	1261.63	2729.23	0.96	2622.24	-
4	14.1	2011	0.85	14.11	3.03	2.41	1464.33	3528.10	0.21	744.08	-
5	18.7	1978	1.27	17.85	5.07	3.38	1552.18	5242.78	0.45	2382.06	-
6	13.7	1958	0.69	13.78	3.03	2.42	1598.54	3867.44	0.24	925.67	-
7	12.8	1988	1.19	12.83	4.26	2.64	731.58	1931.78	0.66	1273.24	-
8	10.4	1975	1.17	10.44	4.62	2.49	1511.67	3758.83	0.51	1929.92	-
9	12.3	2034	1.92	12.36	5.68	2.81	1573.86	4425.98	0.80	3557.16	755.72
10	5.4	2031	0.98	5.38	3.60	1.60	1353.70	2172.14	0.55	1198.15	41.88
11	5.6	2035	0.82	5.70	3.46	1.64	1365.91	2240.67	0.38	855.71	-
12	12.3	2024	2.49	7.50	7.66	2.43	1301.91	3161.48	1.31	4147.94	1511.08
13	11.7	1998	2.02	11.70	5.90	2.77	1416.62	3917.04	0.91	3562.54	2126.86
14	15	1990	1.16	14.97	4.31	2.88	1634.49	4701.79	0.44	2089.65	-
15	13	1991	1.37	12.98	5.72		1584.02	4803.36	0.47	2250.19	
16	13	1960	1.85	13.14	5.58	2.89	1459.94	4219.68	1.08	4566.22	748.13
17	10.8	1938	1.34	10.87	4.74	2.53	1395.21	3532.88	0.54	1902.01	-
18		1940	1.31	15.28	4.50	2.93	1508.62	4425.39	0.55	2422.46	-
19	14.6	1948	1.08	14.69	4.78	3.03	1495.99	4526.71	0.38	1700.80	-
20	14.5	1950	1.40	14.51	4.34	2.79	1492.05	4158.74	0.55	2298.53	117.14
21	11.6	1960	1.89	11.64	6.38	2.90	1420.44	4121.22	0.75	3078.55	-
22	12.7	1992	2.04	12.72	7.47	3.24	1441.28	4665.08	1.06	4954.08	3893.83
23		2013	1.11	12.63	3.96		1578.37	4019.13	0.40	1602.83	120.18
24		2018	1.31	8.56		2.47	1462.07	3608.50	0.56	2028.70	293.83
25		2052	1.21	12.81	5.13	2.89	1587.95	4592.88	0.54	2500.31	22.10
26	5	2030	2.15	4.91	5.80	1.76	1334.68	2343.40	0.83	1946.83	88.99
27		2023	1.02	8.14	3.84	2.03	1449.27	2940.68	0.38	1129.55	116.34
28	9.6	2017	1.45	9.60	5.44	2.52	1493.59	3770.66	0.57	2142.49	443.51

S: the slope of sub-catchments; H: Mean elevation of sub-catchments; L: the length of sub-catchments, D: the difference between mean altitude of the sub-catchments and the altitude of the sub-catchments outlet, P: the perimeter of the sub-catchments, Ru: sediment yield coefficient, Wsp: the average annual erosion, Gsp: Special sediment, A: Area of sub-catchments, Gs: Total sediment: VS: volume of sediment in check dams

Table 4. Fractal dimension and its correlation coefficient (R) for the drainage network in sub-catchments

			` /		
Sub- catchment	FD	R	Sub- catchment	FD	R
1	1.04	0.74	15	1.10	0.83
2	1.1	0.8	16	1.06	0.71
3	1.13	0.79	17	1.09	0.78
4	0.89	0.61	18	1.07	0.79
5	1.11	0.84	19	1.05	0.77
6	1.07	0.62	20	1.09	0.78
7	1.06	0.75	21	1.09	0.77
8	1.14	0.82	22	1.059	0.76
9	1.12	0.84	23	1.121	0.83
10	1.05	0.74	24	1.092	0.79
11	0.93	0.63	25	1.01	0.67
12	1.10	0.78	26	1.08	0.79
13	1.09	0.77	27	1.01	0.66
14	1.11	0.8	28	1.05	0.77

Table 5. The number of sub-catchments located in different sediment transportation zones

	Very High	High	Moderate	Low	Very Low
EPM	5, 15, 14, 22, 25, 19, 9, 18	6, 8, 13, 16, 20, 21, 23, 28	24, 17, 4	12, 27, 3	26, 11, 10, 7, 2, 1
EW	2, 12, 22	3, 7, 9, 13, 14, 16, 18, 20, 21, 23, 26	1, 5, 8, 15, 25, 28	6, 17, 19, 24, 27 4,	
FD	3, 5, 8, 9, 14, 23	2, 12, 13, 15, 17, 20, 21, 24	25, 27	1, 6, 7, 10, 16, 18, 19, 22, 26, 28	4, 11
SL	4	2, 5, 23	3, 7, 9, 12, 16, 18	1, 8, 13, 15, 19, 21, 22, 24, 25, 28	6, 10, 11, 14, 17, 20, 26, 27
sc	6, 18	5, 14, 19, 20, 25	4, 15, 16, 17, 23	7, 8, 9, 13, 21, 22, 24, 27, 28	1, 2, 3, 10, 11, 12, 26

Appendix A Table A1: The SL-based clustering algorithm used in this study

N. C.d.	No. Sub-catchments		1	2	3	4	5	6	7	8	9	10	11	12	13	14
No. Sub-	catenments							Perce	nt of area und	er degree of S	L index					
	o	Very Low	83.48	60.27	85.887	63.317	57.186	73.14	60.227	81.708	60.943	99.281	97.669	86.247	67.476	74.493
	n five	Low	9.756	24.21	7.709	11.943	30.815	24.937	32.549	11.461	30.197	0.7180	2.3042	6.670	26.355	25.011
Stage1	assify in degree	Moderate	4.745	9.440	3.4299	14.834	9.667	1.921	6.191	6.713	6.532	0	0.0261	3.712	5.3106	0.4723
₹	Reclassify in degree	High	1.640	5.172	1.5291	8.720	2.157	0	0.971	0.1167	2.028	0	0	2.302	0.7696	0.0224
	Re	Very High	0.376	0.898	1.444	1.184	0.172	0	0.0606	0	0.298			1.067	0.0879	
	nd	0.04	3.127	2.258	3.217	2.372	2.142	2.740	2.256	3.061	2.283	3.719	3.658	3.231	2.527	2.790
61	hting	0.06	0.537	1.332	0.424	0.657	1.696	1.373	1.792	0.6310	1.6626	0.0395	0.1268	0.3672	1.451	1.377
Stage2	AHP Weighting	0.13	0.611	1.215	0.4417	1.9103	1.2450	0.2474	0.7973	0.8645	0.8412	0	0.0033	0.4780	0.6839	0.0608
Ø		0.25	0.417	1.315	0.388	2.217	0.548	0	0.246	0.029	0.515	0	0	0.585	0.195	0.005
		0.52	0.197	0.471	0.757	0.621	0.090	0	0.031	0	0.156	0	0	0.559	0.046	0
Stage3	Stage3 Specific Number			6.593	5.229	7.778	5.723	4.360	5.124	4.586	5.459	3.758	3.789	5.221	4.9047	4.234
No Sub-	No. Sub-catchments			16	17	18	19	20	21	22	23	24	25	26	27	28
110.040	cutchinonis		Percent of area under degree of SL index													
	,e	Very Low	64.56	59.94	86.812	66.587	78.236	76.853	69.677	74.158	52.733	82.656	61.274	88.324	80.530	70.953
_	in fiy	Low	29.29	34.42	12.667	21.684	16.986	22.097	23.551	20.576	34.653	11.446	35.530	8.654	16.345	24.964
Stage1	ıssify in degree	Moderate	5.909	4.516	0.5014	11.563	4.3939	0.9764	5.9547	4.6030	9.7291	5.1430	3.084	2.3350	2.576	3.6769
N N	Reclassify in five degree	High	0.213	0.997	0.0185	0.1644	0.3538	0.0723	0.7761	0.557	2.557	0.7347	0.1101	0.577	0.4685	0.3870
	ă	Very High	0.021	0.110	0	0	0.0294	0	0.0401	0.1039	0.3259	0.0193	0	0.1083	0.078	0.0175
	ao	0.04	2.418	2.245	3.252	2.494	2.9309	2.8791	2.6103	2.7782	1.975	3.096	2.295	3.3089	3.0169	2.6581
2	ghtin	0.06	1.612	1.895	0.697	1.193	0.935	1.216	1.296	1.132	1.908	0.630	1.956	0.476	0.899	1.374
Stage2	Weig	0.13	0.761	0.581	0.064	1.489	0.565	0.125	0.766	0.592	1.252	0.662	0.397	0.3007	0.331	0.473
N.	AHP Weighting	0.25	0.054	0.253	0.0047	0.0418	0.0899	0.0183	0.197	0.1417	0.6503	0.186	0.028	0.146	0.119	0.0984
	_ <	0.52	0.011	0.058	0	0	0.0154	0	0.021	0.0545	0.1709	0.0101	0	0.0568	0.0409	0.0092
Stage3	Spec	ific Number	4.858	5.034	4.0190	5.2194	4.5375	4.2399	4.8923	4.7002	5.9577	4.5860	4.6770	4.2898	4.4088	4.6138

No. Sub-catchments		1	2	3	4	5	6	7	8	9	10	11	12	13	14	
	. to. buo cutoni		Percent of area under degree of SC index.													
	e ii.	Very Low	8.38	2.31	10.07	0.66	3.98	0.46	5.98	0.84	1.89	2.47	17.83	12.88	6.63	0.85
gel	ssify	Low	13.46	13.37	39.51	1.99	3.83	1.38	13.44	9.44	9.85	30.20	22.70	16.70	13.60	5.04
Sta	eclas	Moderate	36.83	38.97	26.81	15.36	13.76	8.98	21.05	25.10	27.16	27.97	25.22	27.75	27.07	14.93
	R. B.	High	26.43	31.01	12.15	49.62	35.11	26.07	23.81	36.49	39.34	30.57	21.37	36.44	29.75	37.25

Appendix B Table B1 The algorithm of clustering using SC index

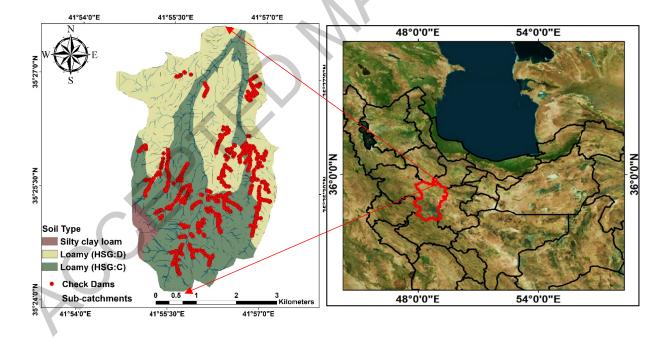
		Very High	14.9	14.34	11.46	32.37	43.32	63.12	35.72	28.14	21.75	8.79	12.88	6.23	22.95	41.93
		0.04	0.31	0.09	0.38	0.02	0.15	0.02	0.22	0.03	0.07	0.092	0.668	0.483	0.248	0.032
	g	0.04	0.74	0.74	2.18	0.02	0.13	0.02	0.22	0.52	0.54	1.663	1.25	0.483	0.749	0.032
e2	AHP Weighting	0.00	4.74	5.02	3.45	1.98	1.77	1.16	2.71	3.23	3.5	3.602	3.247	3.573	3.486	1.923
Stage 2	ĕ	0.13	6.72	7.88	3.43	12.6	8.93	6.63	6.05	9.28	10	7.772	5.433	9.266	7.565	9.47
	AHI	0.23	7.81	7.52	6.01	17.0	22.7	33.1	18.7	14.8	11.4	4.611	6.756	3.266	12.03	21.99
Ct. 2	C						11									
Stage3	Stage3 Specific Number		20.3	21.2	15.1	31.7	33.8 19	41 20	28.5	27.8	25.5	17.74	17.35 25	17.51 26	24.08	33.69
l l	No. Sub-catchi	ments	15	16	17	18	19	-	21		23	24	25	26	27	28
		T	Percent of area under degree of SC index. 1.56													
	.se	Very Low														
20	in f	Low	6.06	5.65	3.71	1.52	0.69	3.71	23.23	11.43	7.15	17.31	3.12	35.26	6.09	10.97
Stage1	assify in degree	Moderate	20.5	17.82	18.26	7.73	8.89	14.61	20.58	28.12	13.47	25.79	13.35	25.71	28.72	21.25
01	eclas	High	38.83	41.35	48.70	31.95	48.12	28.33	22.76	37.03	40.15	37.32	41.71	17.56	43.67	39.33
	N.	Very High	33.05	34.14	29.22	58.75	42.16	52.70	31.23	20.50	37.11	17.72	41.76	2.88	20.36	20.32
	50	0.04	0.058	0.039	0.004	0.002	0.005	0.024	0.082	0.109	0.08	0.07	0.002	0.697	0.043	0.305
7	ghtin	0.06	0.334	0.311	0.205	0.083	0.038	0.204	1.279	0.63	0.393	0.953	0.172	1.941	0.336	0.604
Stage 2	Wei	0.13	2.639	2.295	2.351	0.995	1.145	1.881	2.65	3.621	1.734	3.321	1.719	3.311	3.699	2.737
σ.	AHP Weighting	0.25	9.874	10.51	12.38	8.124	12.23	7.203	5.787	9.416	10.21	9.488	10.6	4.465	11.1	10
	- ■	0.52	17.33	17.9	15.32	30.81	22.11	27.64	16.38	10.75	19.46	9.291	21.9	1.509	10.68	10.65
Stage3	Spec	eific Number	30.24	31.06	30.26	40.02	35.53	36.95	26.18	24.53	31.88	23.12	34.4	11.92	25.86	24.3
				PC C		R										

Figure captions

- **Figure 1.** Location of the Ilanlu catchment in Hamadan Province and Iran and examples of check dams with collecting sediments
- Figure 2. Flowchart of the proposed methods
- Figure. 3. An example of calculating the fractal dimension in sub-catchments
- Figure. 4. Mapping of sediment production hotspots using FD (a), EW (b), SL (c) and SC (d) methods
- **Figure. 5.** Clustering similar sub-catchments in terms of sediment production based on FD (a), EW (b), SL (c) and SC (d) results
- **Figure. 6**. above) Schematic of a box-counting method to evaluate the presence of a drainage network at different sizes and below) A normalized evaluation of the presence of a river network for a variety of different sizes and calculating their entropy

Figure. 7. above) The SL and SC values at sub- catchments and below) The standard deviation (SD) of SL and SC indices at sub-catchments

Figure 8. above) Mapping of sediment production hotspots and below) Clustering similar sub-catchments in terms of sediment production using EPM model



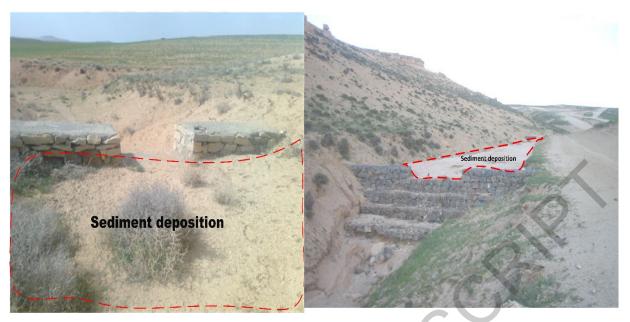


Figure 1. Location of the Ilanlu catchment in Hamadan Province and Iran and examples of check dams with collecting sediments

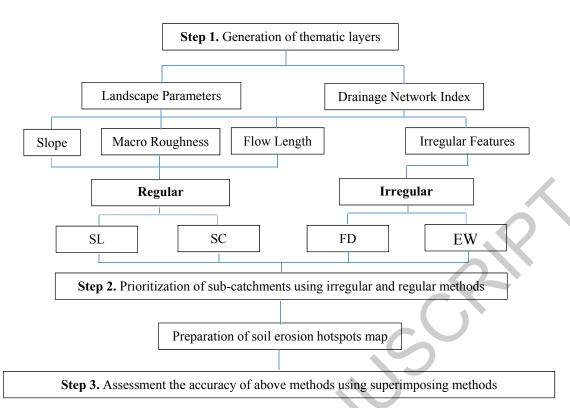


Figure 2. Flowchart of the proposed methods

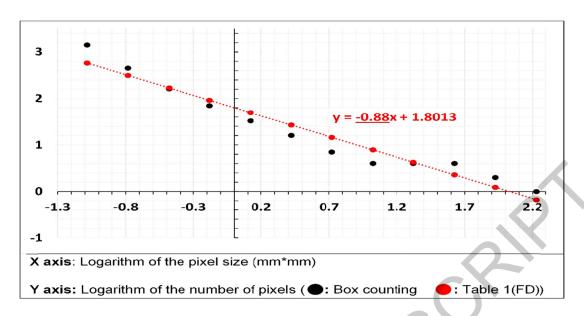


Fig. 3. An example of calculating the fractal dimension in sub-catchments

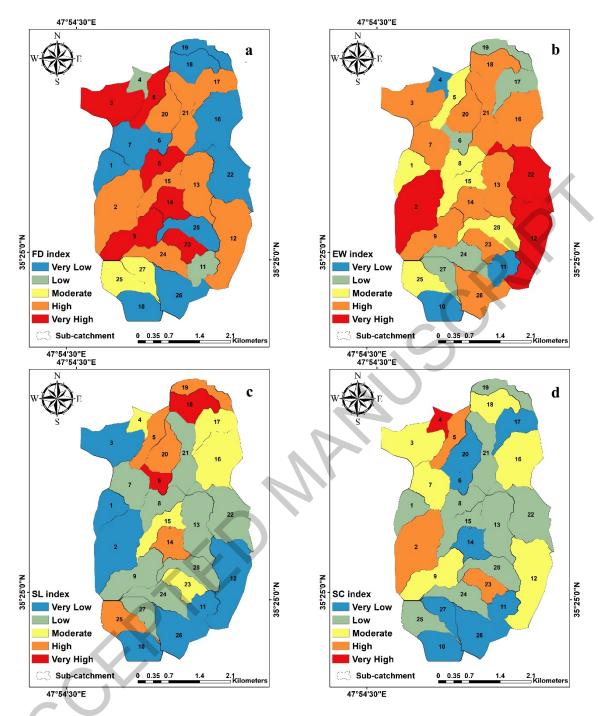


Fig. 4. Mapping of sediment production hotspots using FD (a), EW (b), SL (c) and SC (d) methods

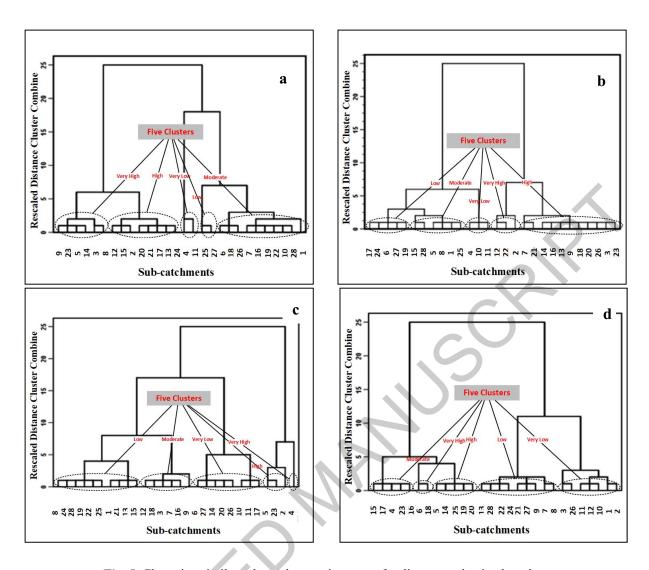


Fig. 5. Clustering similar sub-catchments in terms of sediment production based on FD (a), EW (b), SL (c) and SC (d) results

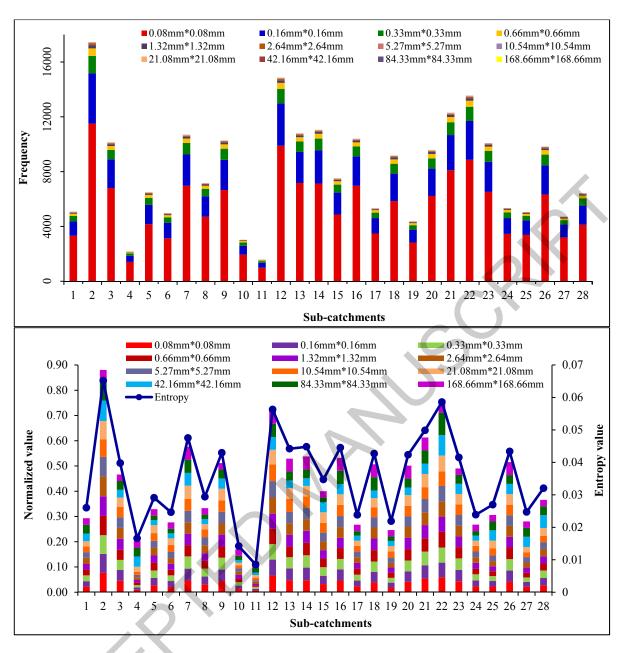


Fig.6. above) Schematic of a box-counting method to evaluate the presence of a drainage network at different sizes and below) A normalized evaluation of the presence of a river network for a variety of different sizes and calculating their entropy

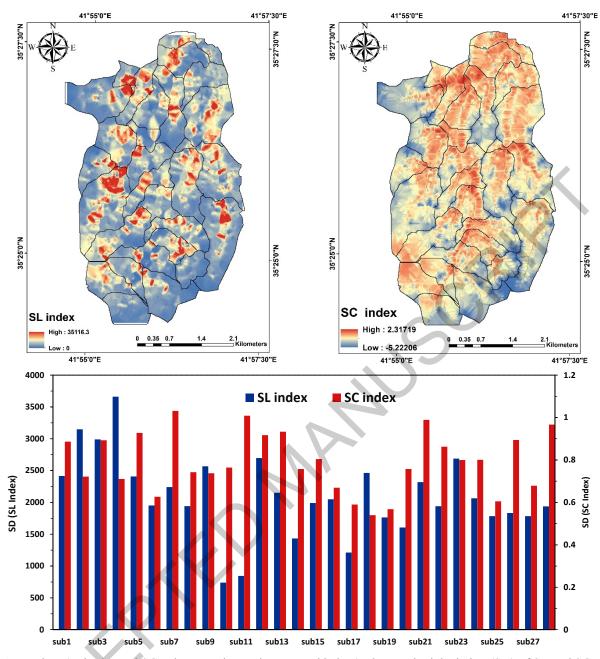


Fig. 7. above) The SL and SC values at sub- catchments and below) The standard deviation (SD) of SL and SC indices at sub-catchments

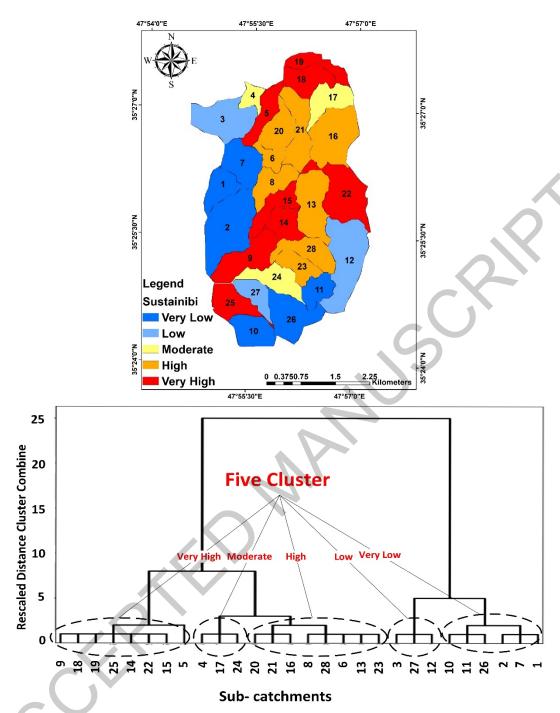


Fig 8. above) Mapping of sediment production hotspots and below) Clustering similar sub-catchments in terms of sediment production using EPM model